DIAMONDS AND MANTLE XENOLITHS IN KIMBERLITES FROM THE NORTH CHINA CRATON AND THE CANADIAN NORTHWEST TERRITORIES

by

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Dedicated to my families, in China and here in the United States, especially to my wife Qinglan Liu and children Sijia and Sihua.

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TABLE OF CONTENTS

DED	ICA	TION	ii
ACKN	NOW	LEDGMENTS	iii
LIST	OF	FIGURES	.viii
LIST	OF 1	ΓABLES	.xvi
LIST	OF	APPENDICES	xviii
ABST	RAC	CT xix	
СНАР	PTEF	R	
	I.	INTRODUCTION	1
	II.	MINERAL INCLUSIONS IN DIAMONDS FROM THE NO. 50 KIMBERLITE DIATREME, LIAONING PROVINCE, CHINA	4
		Abstract. Introduction Geology of the No. 50 kimberlite diatreme. Diamonds studied and their morphology Sample preparation. Analytical methods Features of mineral inclusions and their hosts Chemistry of mineral inclusions. Discussion. Conclusions. References cited.	4 5 11 16 17 18 41 77 89 91
	III.	MINERAL INCLUSIONS IN CHROMITES FROM THE FUXIAN KIMBERLITES, LIAONING PROVINCE, CHINA	99

Abstract	99
	<i>'</i> -

	Introduction Geology of the No. 42 and 50 kimberlite diatremes Analytical methods Features of chromites and their mineral inclusions Compositions of chromites and their mineral inclusions Discussion References cited	100 100 101 103 115 136 145
IV.	MANTLE XENOLITHS FROM THE NIKOS KIMBERLITES ON SOMERSET ISLAND AND THE ZULU KIMBERLITES ON BRODEUR PENINSULA, BAFFIN ISLAND, CANADA	. 148
	Abstract Introduction The Nikos and Zulu kimberlites Petrography Mineral chemistry P-T estimates of xenoliths Oxygen fugacities Discussion References.	148 149 151 152 155 172 174 174 181
v.	AN OXYGEN BAROMETER FOR RUTILE-ILMENITE ASSEMBLAGES: OXIDATION STATE OF METASOMATIC AGENTS IN THE MANTLE	. 186
	Abstract Introduction Formulation of the oxygen barometer Sources of error Comparison with other oxygen barometers Applications References.	186 186 189 191 196 196 201
VI.	CONCLUSIONS	206
APPEND	DICES	209

LIST OF FIGURES

Figures of Chapter II

Fig. 2.1.	Schematic tectonic map of China showing the main tectonic elements and kimberlite and lamproite localities in the North China Craton
Fig. 2.2.	The open pit of the No. 50 kimberlite diatreme, Liaoning Province, China (photo by Donggao Zhao, June, 1995). The field of the photo is about 300 meters across
Fig. 2.3.	Schematic diagrams of the No. 50 kimberlite diatreme outcrop (A) and profile (B), Liaoning Province, China10
Fig. 2.4.	Diamond crystals: (A) octahedral diamond (No. 4), size ~2 mm; (B) dodecahedral diamond (No. 3), size ~3 mm; (C) octahedral diamond (No. 14), size ~1 mm. From Huang et al. (1992)15
Fig. 2.5.	Transmitted light image of diamond LN50D02 with two big and one small olivine inclusion. Notice the black platelets evenly distributed over the surface of the small euhedral elongated olivine inclusion. Plane polarized light. Width of field ~1.2 mm
Fig. 2.6.	Transmitted light images of two oriented olivine inclusions (A) and an irregular chromite inclusion (B) in diamond LN50D04. Plane polarized light. Width of field: (A) ~ 0.2 mm; (B) ~ 0.1 mm20
Fig. 2.7.	Transmitted light image of diamond LN50D07 with two chromite inclusions. One of the inclusions is only partly included in the diamond host. Plane polarized light. Width of field ~2 mm
Fig. 2.8.	Transmitted light image of diamond LN50D10 with garnet (purple) and olivine (lower right corner) inclusions. Plane polarized light. Width of field ~1.5 mm
Fig. 2.9.	Transmitted light image of diamond LN50D40 with two orthopyroxene inclusions. Plane polarized light. Width of field ~0.4 mm21
Fig. 2.10.	Diamond LN50D38 with butterfly-like graphite inclusions. The X-ray mapping of the fracture shows no Si counts, thus indicating that the inclusions are not SiC. Transmitted light, plane polarized light. Width of field: (A) ~ 2 mm; (B) ~ 0.3 mm
Fig. 2.11.	Multiple inclusions: (A) Diamond LN50D12 with at least 7 inclusions (4 olivines and 3 chromites); (B) Diamond LN50D14 with at least 6 inclusions (most are olivines). Transmitted light, plane polarized light. Width of field ~2 mm

- Fig. 2.12. BSE images of two polished sections for diamond LN50D03 (A) with an elongated olivine inclusion, showing cathodoluminescent patterns. Width of field ~1.3 mm; (B) two euhedral olivine inclusions. The c axis of the large olivine is perpendicular to the polished section, while the c axis of the small olivine is close to parallel to the polished section. Width of field ~1.1 mm. Working distance 15 mm, accelerating voltage 15 kV.......24
- Fig. 2.13. BSE images of polished section of diamond LN50D04 (A) two olivine inclusions and one chromite inclusion at the upright corner. Notice the cathodoluminescent pattern surrounding the olivine at the center. Width of field ~1.7 mm; (B) an euhedral olivine inclusion at the center of the diamond. Notice the ring structure surrounding the olivine inclusion, which may be the result of strain between the inclusion and host. Width of field 0.1 mm. Working distance 15 mm, accelerating voltage 15 kV.......24
- Fig. 2.14. BSE image of polished section of diamond LN50D39 with an olivine inclusion, showing triangular cathodoluminescent pattern (also see LN50D40). Width of field ~2.0 mm. Working distance 14 mm, accelerating voltage 20 kV.....25
- Fig. 2.16. BSE images of polished section of diamond LN50D45 (A) the host with an euhedral olivine inclusion. Width of field 2.1 mm; (B) the euhedral olivine inclusion at higher magnification. Notice the bent feature of the crystal, which was constrained by the crystal form of diamond host. Width of field 0.2 mm. Working distance 15 mm, accelerating voltage 20 kV.......26

- Fig. 2.21. BSE images of polished section of diamond LN50D12 (A) the host with a chromite inclusions at the center of cathodoluminescent pattern. Width of

field 2.0 mm; (B) the chromite inclusion at higher magnification. Width of field 0.18 mm. Working distance 15 mm, accelerating voltage 15 kV......28

- Fig. 2.25. BSE images of polished section of the diamond LN50D71 (A) the host with a garnet inclusion. Notice the complex cathodoluminescent pattern. Width of field ~1.5 mm; (B) the garnet inclusion at higher magnification. Width of field 0.06 mm. Working distance 15 mm, accelerating voltage 15 kV......31

- Fig. 2.32. (A) BSE image of polished section of diamond LN50D45 with a chromite inclusion in the broken fracture, which was developed during polishing. Width of field 1.8 mm; (B) secondary electron image of the chromite inclusion in diamond LN50D45 at higher magnification. Width of field 0.2 mm. Working distance 15 mm, accelerating voltage 20 kV......37
- Fig. 2.33. Sulfide in diamond LN50D04. Fe and Ni are heterogeneously distributed and Cu is much enriched in some areas. The view area is ~30 µm......38

- Fig. 2.41. Diagrams of CaO versus Cr_2O_3 (A), Mg # versus Cr_2O_3 (B), and Mg # versus TiO_2 (C) for garnet inclusions in diamonds. Solid line divides

compositional fields for lherzolitic and harzburgitic garnets, dashed line distinguishes peridotitic from non-peridotitic garnets. Both lines are described in detail in Gurney and Zweistra (1995)......57

- Fig. 2.42. Histograms of Mg# (A), Cr₂O₃ (B) and SiO₂ (C) for chromite inclusions in diamonds. The data are listed in Table 2.7......62
- Fig. 2.43. Diagrams of Cr₂O₃ versus TiO₂ (A) and Fe₂O₃ versus Al₂O₃ (B) for chromite inclusions in diamonds from the No. 50 kimberlites. Dashed lines delineate a "diamond inclusion field" (Kopylova et al. 1997)63

Figures of Chapter III

Fig. 3.1.	Fragment of an anhedral chromite host (LN42SP10) with a hole left by inclusion
Fig. 3.2.	An anhedral chromite host (LN42SP12) with a small inclusion104
Fig. 3.3.	Rounded, anhedral chromite (A) with an euhedral olivine inclusion (B). Sample # LN42SP14105
Fig. 3.4.	Chromite host (LN50SP01) with a silicate inclusion, which has magnetite crystals on the rim. The original inclusion might have been orthopyroxene or olivine
Fig. 3.5.	Chromite host (LN50SP04) with olivine inclusion106
Fig. 3.6.	Subhedral chromite (A) with an euhedral, elongated olivine inclusion (B). Sample # LN42SP13106
Fig. 3.7.	Chromite host LN42SP01(A) with two olivine inclusions (B and C) and one pyrope (D). The garnet inclusion possesses a symplectitic rim in which aluminous spinel and pyroxene were identified. (E) and (F) are enlarged images of squares in (D)
Fig. 3.8.	Subhedral chromite (A) with a subhedral olivine inclusion (B). Sample # LN42SP16108
Fig. 3.9.	Chromite host (LN50SP02) with olivine inclusion108
Fig. 3.10.	Chromite host (LN50SP03) with olivine inclusion109

Fig. 3.11.	Chromite host (LN50SP05) with a magnesite inclusion (B). The longer bar in (B) is 10 µm109
Fig. 3.12.	Chromite host (LN50SP06) with olivine inclusion110
Fig. 3.13.	Euhedral chromite with euhedral olivine inclusion (LN42SP2)110
Fig. 3.14.	Euhedral chromite (A) with a hole (B) left by the inclusion (LN42SP04)111
Fig. 3.15.	Chromite host (LN50SP10) with a sulfide inclusion111
Fig. 3.16.	A rounded, anhedral olivine inclusion in chromite LN42SP11112
Fig. 3.17.	Chromite host (LN42SP07) with phlogopite (A), calcite (B), and a complicated inclusion of norsethite and phlogopite (C)112
Fig. 3.18.	Chromite host (LN42SP06) with a single composite inclusion containing dolomite, norsethite and phlogopite113
Fig. 3.19.	Chromite host (LN50SP07) with) with a single composite inclusion containing magnesite, dolomite, norsethite, phlogopite and SiO_2 113
Fig. 3.20.	Histograms of Mg# (A), Cr_2O_3 (B) and SiO ₂ (C) for chromites in the No 42 and 50 kimberlite diatremes, Liaoning. The data are listed in Table 3.2118
Fig. 3.21.	Diagrams of Cr_2O_3 versus TiO_2 (A) and Fe_2O_3 versus Al_2O_3 (B) for chromites from the Liaoning kimberlites. Dashed lines in (A) delineate a "diamond inclusion field" (Kopylova et al. 1997). Dotted lines in (B) delineate a diamond inclusion field for the No. 50 kimberlites
Fig. 3.22.	Histograms of Mg# (=100Mg/(Mg+Fe) by atoms) (A) and NiO content (B) for the olivine inclusions in chromites from the No. 42 and 50 kimberlite diatremes
Fig. 3.23.	Mg # versus CaO (wt %) for the olivine inclusions in chromites from the No. 42 and 50 kimberlite diatremes
Fig. 3.24.	Variation of Cr across an olivine inclusion in chromite LN50SP06 (961008, 10 kV; 961010, 30 kV). The increase of Cr toward the chromite-olivine boundary is owing to fluorescence. Cr fluorescent effect becomes significant with increasing voltage or approaching the olivine-chromite boundary124
Fig. 3.25.	Fe ²⁺ variation across an olivine in chromite LN42SP01138
Fig. 3.26.	Temperature-pressure plots of each assemblages for chromites from the No. 42 and 50 kimberlite diatremes. Each curve represents a thermobarometer. See text and Table 3.10 for abbreviations

Figures of Chapter IV

Fig. 4.1. Map of Somerset Island and Brodeur Peninsula, Baffin Island, Northwest Territories, Canada, showing kimberlite localities......150

- Fig. 4.4. CaO vs. Cr₂O₃ diagram for garnet from the Nikos and Zulu xenoliths165

Figures of Chapter V

- Fig. 5.1. logfO₂ vs. *T* for several oxygen buffers at 1 bar. MH: magnetite-hematite;
 RI: rutile-ilmenite; NNO: Ni-NiO; QFM: quartz-fayalite-magnetite; WM: wüstite(Fe_{0.947}O)-magnetite. The RI buffer curves are calculated using data from [28-30] for comparison. The thick curve of RI shows overlapping from [29] and [30]; the thin curve below is from [28] with modification from Ghiorso [26]. Other buffers are calculated using data from [30] only.....190
- Fig. 5.2. Comparison of calculated Fe³⁺/(Fe³⁺+Fe²⁺) ratios using microprobe analyses from various laboratories (vertical axis) with the Mössbauer measurements of Virgo et al. [10]. Three samples (ULM 2, Yakutia-Dalnaya No. 1 D-46/79, and ROM 264 IL-41) used in the comparison were provided by D. Virgo. Legend: UM = University of Michigan, ZAF correction; PSU = Pennsylvania State University, Bence-Albee correction; GL-BA=Geophysical Lab, Bence-Albee correction; GL-ZAF=Geophysical Lab, ZAF correction. Results of UM microprobe analyses are from this work. Results of other microprobe

analyses are from Virgo et al. [10]. The solid line is a 1:1 line. Also shown are 2 error bars for Mössbauer analyses (3% relative, D. Virgo, pers. comm.) and for repeated microprobe analyses at University of Michigan ...195

LIST OF TABLES

Tables of Chapter II

Table 2.1.	Diamonds and their mineral inclusions from the No 50 kimberlite diatreme, Liaoning, China
Table 2.2.	Mineral inclusions identified in diamonds from the No. 50 kimberlite diatreme
Table 2.3.	Average compositions of olivine inclusions in diamonds from the No. 50 kimberlite diatreme
Table 2.4.	Compositions of orthopyroxene inclusions in diamonds from the No. 50 kimberlite diatreme
Table 2.5.	Compositions of clinopyroxene inclusions in diamonds from the No. 50 kimberlite diatreme
Table 2.6.	Compositions of garnet inclusions in diamonds from the No. 50 kimberlite diatreme
Table 2.7.	Compositions of chromite inclusions in diamonds from the No. 50 kimberlite diatreme
Table 2.8.	Representative compositions of sulfide inclusions in diamonds from the No. 50 kimberlite diatreme
Table 2.9.	Compositions of carbonate inclusions in diamonds from the No. 50 kimberlite diatreme71
Table 2.10.	Compositions of unknown (dense hydrous?) phases in diamonds from the No. 50 kimberlite diatreme
Table 2.11.	Compositions of unknown Fe-rich phase (goethite?) in diamond from the No. 50 kimberlite diatreme
Table 2.12.	Calculated T (°C) and P (kbar) for mineral inclusion assemblages in diamonds from the No. 50 kimberlites

Tables of Chapter III

Table 3.1.	Chromites and their mineral inclusions from the No 42 and 50 kimberlit	e
	diatremes, Liaoning11	4

Table 3.2.	Compositions of chromites from the No. 42 and 50 diatremes116
Table 3.3.	Compositions of olivine inclusions in chromites from the No. 42 and 50 kimberlite diatremes
Table 3.4.	Compositions of a garnet inclusion in chromite LN42SP01126
Table 3.5.	Compositions of pyroxenes in the symplectite rim of a garnet inclusion in chromite LN42SP01127
Table 3.6.	Compositions of phlogopite inclusions in chromite LN42SP07129
Table 3.7.	Compositions of an unknown silicate inclusion in chromite LN50SP01 .130
Table 3.8.	Compositions of carbonate inclusions in chromites from the No. 42 and 50 kimberlite diatremes
Table 3.9.	Compositions of a sulfide inclusion in chromite LN50-SP10135
Table 3.10.	Calculated T (°C) and P (kbar) for mineral inclusions in chromites from the Liaoning kimberlites
Table 3.11.	Comparison of mineral inclusions in diamonds and in chromites from the Liaoning kimberlites

Tables of Chapter IV

Petrography, pressure, temperature and oxygen fugacity of the mantle- derived xenoliths from the Nikos and Zulu kimberlites153
Average compositions of olivines in the Nikos and Zulu xenoliths158
Average compositions of Opx in the Nikos and Zulu xenoliths160
Average compositions of Cpx in the Nikos and Zulu xenoliths161
Average compositions of garnets in the Nikos and Zulu xenoliths164
Average compositions of spinels in the Nikos and Zulu xenoliths167
Average compositions of phlogopites in the Nikos and Zulu xenoliths and kimberlites
Compositions of ilmenites and rutiles in the Nikos xenoliths170
Compositional profile across an ilmenite (~300 µm) adjacent to the spinel in the JP1-X17 MORID vein

Table of Chapter V

Table 5.1.	Oxygen fugacities of	various	assemblages	calculated	from	rutile-ilmenite
	(RI) and other oxyg	en baro	meters		•••••	197

LIST OF APPENDICES

Appendices of Chapter II

Appendix	2.1. Compositions of olivine inclusions in diamonds from the No. 50 kimberlite diatreme
Appendix	2.2. Compositions of orthopyroxene inclusions in diamonds from the No. 50 kimberlite diatreme
Appendix	2.3. Compositions of garnet inclusions in diamonds from the No. 50 kimberlite diatreme
Appendix	2.4. Compositions of chromite inclusions in diamonds from the No. 50 kimberlite diatreme
Appendix	2.5. Compositions of sulfide inclusions in diamonds from the No. 50 kimberlite diatreme
Appendix	2.6. Average compositions of peridotitic mineral inclusions in diamonds from the Liaoning kimberlites

Appendices of Chapter III

Appendix 3.1. Compositions of chromites from the N	Jo. 42 and No. 50 kimberlite
diatremes	
Appendix 3.2. Compositions of olivine inclusions in chr	omites from the No. 42 and No.
50 kimberlite diatremes	

Appendix of Chapter IV

Appendix 4.1.	Microprobe	analyses of	of the	mantle-der	rived 2	xenoliths	from	the 1	Nikos	and
Zul	u kimberlites	by sample	e					••••		311

ABSTRACT

This dissertation focuses on mantle compositions and processes. Mineral inclusions in diamonds from the Liaoning kimberlites, China belong mainly to a harzburgitic assemblage. The diamonds crystallized at depths of 140 to 200 km. Mineral inclusions in chromites from the same locality include silicates, carbonates, hydrous silicates and sulfides. Composite inclusions of carbonates+silicates in chromites might represent entrapped and crystallized kimberlitic magma. A garnet-olivine-chromite assemblage indicates that the depth where chromites formed is ~140 km. Similarities and differences between the inclusion assemblages in diamonds and in chromites suggest that minerals in diamonds and in chromites came from different depths, and that most mineral inclusions in chromites were probably trapped during the stage of the formation of kimberlite.

Mantle xenoliths from the Nikos kimberlites, Somerset Island, and the Zulu kimberlites, Brodeur Peninsula, Baffin Island, Canada are mainly coarse, protogranular, low-temperature lherzolites. High-temperature xenoliths, which are common in the Kaapvaal and Siberian cratons, were not found at Nikos. Garnet-spinel lherzolite xenoliths are common at Nikos. The calculated pressures and temperatures follow a continental geotherm. The fO_2 from olivine-orthopyroxene-spinel is from 1.3 log units above to 0.6 log units below EMOD (enstatite-magnesite-olivine-diamond), suggesting that diamond may or may not be stable relative to carbonates.

A MORID vein (mica-orthopyroxene-rutile-ilmenite-diopside±chromite) in a garnetspinel lherzolite is characterized by high K, Fe, Ti and OH components. A method (referred to as RI) is developed to calculate oxygen fugacity from rutile-ilmenite for a MORID or similar suite with the reaction $2Fe_2O_3$ (in ilmenite) + $4TiO_2$ (rutile) = $4FeTiO_3$ (in ilmenite) + O_2 . The RI is applicable to many rutile-ilmenite-bearing assemblages.